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12.510 Introduction to Seismology
Spring 2008

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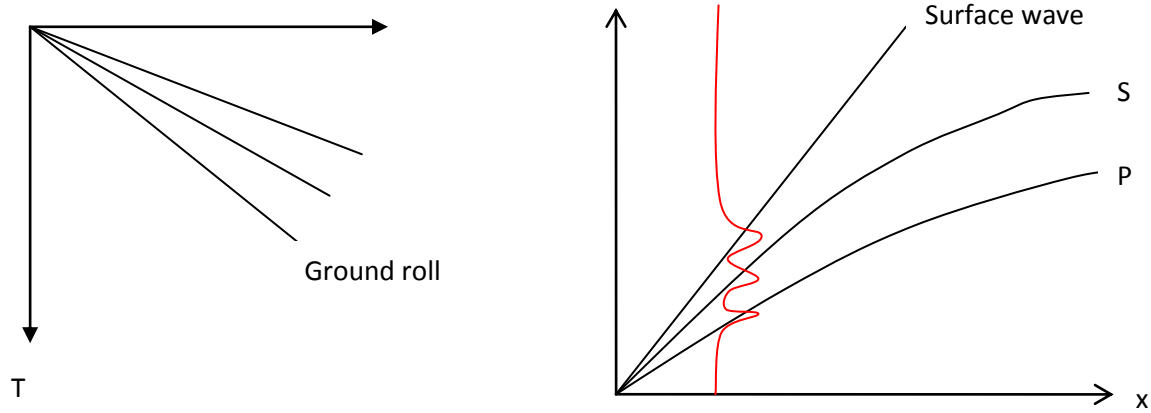
Continents: Quick review. Surface waves

Ground roll-acoustic $\ddot{p} = V \cdot \nabla \cdot \left(\frac{1}{\rho} \nabla p \right)$, where p is the pressure

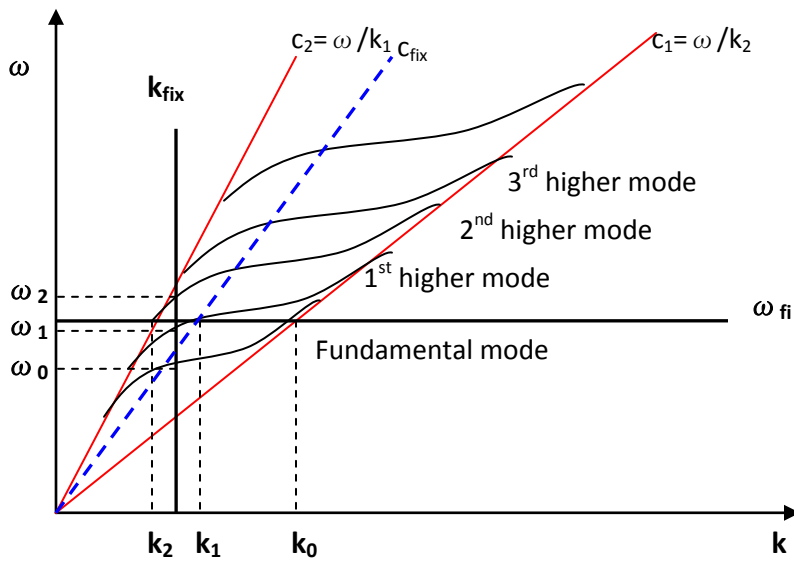
Love waves – SH $\ddot{u}_y = \mu \cdot \nabla \cdot \left(\frac{1}{\rho} \nabla u_y \right)$

Rayleigh waves P-SV

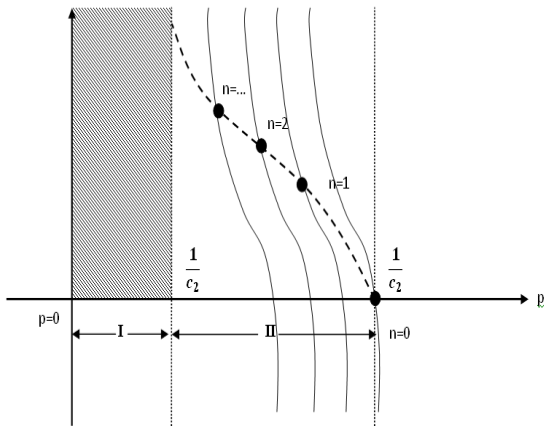
Quick review (refer to April, 4, 2008 for details)



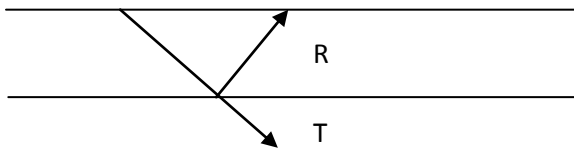
$\omega - k$ domain



Ground roll dispersion relationship

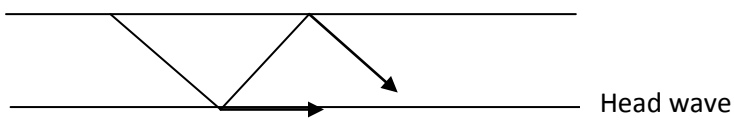


Pre-critical $p < \frac{1}{c_2}$

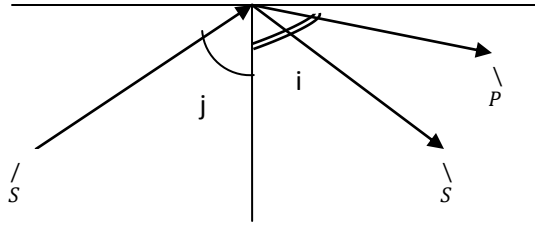


$$\eta_2 = \sqrt{\frac{1}{c_2^2} - p^2} \in \mathbb{R} \tag{1}$$

Post-critical $p > \frac{1}{c_2}$

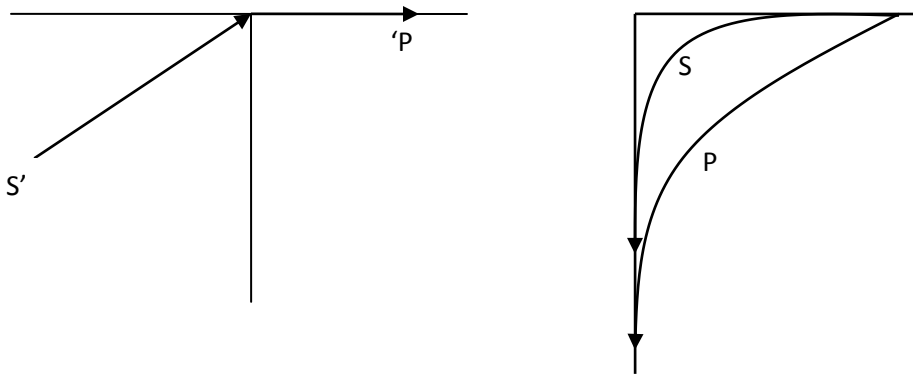


$$\eta_2 = i \sqrt{p^2 - \frac{1}{c_2^2}} = i\eta_2 \in \mathbb{C} \tag{2}$$



$$\frac{\sin(j)}{\beta} = \frac{\sin(i)}{\alpha} = p \tag{3}$$

if $\beta < \alpha$, there can be critical reflection and horizontal propagating p-wave. If $j > j_c$ then there will be evanescence in the p-wave ($p > 1/\alpha$).



Both p-wave and s-wave horizontal propagation if $p > \frac{1}{\beta} > \frac{1}{\alpha} = \frac{1}{c}$. If a wave comes in with a $1/c$ that is larger than local $1/\alpha$ and $1/\beta$, the above will occur. This will also happen if the source emits a horizontal energy (rare).

$$P: \phi = A \exp(-\omega \eta_\alpha z) \exp(i\omega(px - t)), \eta_\alpha = \sqrt{\frac{1}{\alpha^2} - p^2} \underset{p > \frac{1}{\alpha}}{=} i \sqrt{p^2 - \frac{1}{\alpha^2}} = i \widehat{\eta}_\alpha = i \sqrt{\left(\frac{1}{c}\right)^2 - \left(\frac{1}{\alpha}\right)^2}; \text{ where } c = c_R = \frac{1}{p} \text{ (} c_R \text{: phase velocity for Rayleigh waves)} \tag{4}$$

$$S: \psi = \beta \exp(-\omega \eta_\beta z) \exp(i\omega(px - t)), \eta_\beta = \sqrt{\frac{1}{\beta^2} - p^2} \underset{p > \frac{1}{\beta}}{=} i \sqrt{p^2 - \frac{1}{\beta^2}} = i \widehat{\eta}_\beta = i \sqrt{\left(\frac{1}{c}\right)^2 - \left(\frac{1}{\beta}\right)^2} \tag{5}$$

We follow the same “Recipe” we used before:

1. Potentials
2. Boundary Conditions (Kinematic and dynamic)
3. Zoeppritz equations.

Boundary conditions

$$u(x, t) = \nabla\phi + \nabla \times \psi \quad (6)$$

In this case

$$R_{ss}, R_{sp}, R_{ps}, R_{pp} \quad (7)$$

After some work we get:

$$\begin{aligned} A[(\lambda + 2\mu)\eta_\alpha^2 + \lambda p^2] + 2\mu p \eta_\beta &= 0 \\ A(2p\eta_\alpha) + B(p^2 - \eta_\beta^2) &= 0 \end{aligned} \quad (9)$$

Zoeppritz

$$\begin{bmatrix} (\lambda + 2\mu)\eta_\alpha^2 + \lambda p^2 & 2\mu p \eta_\beta \\ 2p\eta_\alpha & p^2 - \eta_\beta^2 \end{bmatrix} \begin{bmatrix} A \\ B \end{bmatrix} = \begin{bmatrix} 0 \\ 0 \end{bmatrix} \quad (10)$$

Trivial solution is

$$\mathbf{A} = \mathbf{B} = \mathbf{0}$$

Non-trivial solution leads to:

$$[(\lambda + 2\mu)\eta_\alpha^2 + \lambda p^2][p^2 - \eta_\beta^2] - 2p\eta_\alpha(2\mu p \eta_\beta) = 0 \quad (11)$$

This expression; Rayleigh wave denominator (Rayleigh, 1887), can be written looking at wave speeds, but usually done numerically assuming a Poisson’s medium ($\lambda=\mu$) and $\alpha = \sqrt{3}\beta$. This scaling will help to simplify the above equation.

$$\begin{aligned} \eta_\alpha &= \sqrt{\frac{1}{\alpha^2} - p^2} = \sqrt{\frac{1}{\alpha^2} - \left(\frac{1}{c}\right)^2} \\ \eta_\beta &= \sqrt{\frac{1}{\beta^2} - p^2} = \sqrt{\frac{1}{\beta^2} - \left(\frac{1}{c}\right)^2} \end{aligned} \quad (12)$$

$$\alpha = \sqrt{\frac{\lambda + 2\mu}{\rho}}$$

$$\Leftrightarrow \rho\alpha^2 = \lambda + 2\mu$$

$$\beta = \sqrt{\frac{\mu}{\rho}}$$

$$\Leftrightarrow \rho\beta^2 = \mu$$

$$\left[\alpha^2 \left(\frac{\eta_\alpha^2}{\rho^2} + 1 \right) - 2\beta^2 \right] \left(1 - \frac{\eta_\beta^2}{\rho^2} \right) - \left(\frac{4\beta^2 \eta_\alpha \eta_\beta}{\rho^2} \right) = 0$$

$$\Leftrightarrow \frac{\beta}{\alpha}$$

Assumptions

Poisson medium: $\lambda = \mu$

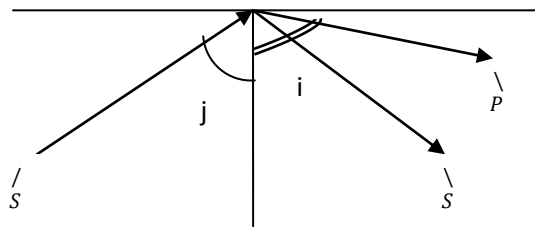
Poisson ratio: $V = \frac{\lambda}{2(\lambda + \mu)} = 0.25$

$$\alpha = \sqrt{3}\beta$$

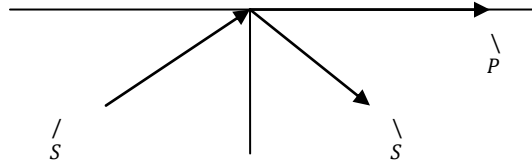
$$\left(\frac{c_R^2}{\beta^2} \right)^3 - 8 \left(\frac{c_R^2}{\beta^2} \right)^2 + \frac{56}{3} \left(\frac{c_R^2}{\beta^2} \right) - \frac{32}{3} = 0$$

Three solutions

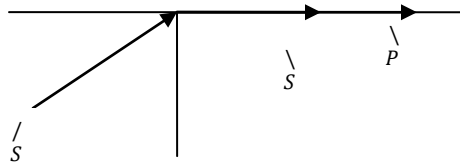
$$1. \left(\frac{c}{\beta} \right)^2 = 4 \rightarrow c = 2\beta, c > \beta, p = \frac{1}{c} < \frac{1}{\beta}, p < \frac{1}{\alpha}$$



$$2. \left(\frac{c}{\beta}\right)^2 = 2 + \frac{2}{3}\sqrt{3}, \frac{1}{\alpha} < p < \frac{1}{\beta}, c^2 = \left(2 + \frac{2}{3}\sqrt{3}\right)\beta^2$$



$$3. \left(\frac{c}{\beta}\right)^2 = \left(2 - \frac{2}{3}\sqrt{3}\right), \frac{1}{\alpha} < p < \frac{1}{\beta}, c = \sqrt{2 + \frac{2}{3}\sqrt{3}}, c < \beta, p = \frac{1}{c} > \frac{1}{\beta} > \frac{1}{\alpha}, (\eta_\alpha = i\widehat{\eta}_\alpha, \eta_\beta = i\widehat{\eta}_\beta)$$



Note: $c_R \sim 0.92\beta$ which makes it about 10% slower than the shear wave velocity. This explains why the Rayleigh wave is slower than the love wave. The Love wave, at low frequency $\rightarrow c_2$ and at high frequencies Love wave (at it's slowest) is c_1 ...Rayleigh wave is about 90% of the Love wave at the most.

exponential decay with depth

$$\begin{aligned} \phi &= A \exp(-\omega \widehat{\eta}_\alpha z) \exp(i\omega(px - t)) \\ \psi &= \beta \exp(-\omega \widehat{\eta}_\beta z) \exp(i\omega(px - t)) \end{aligned} \quad (13)$$

$$u = \nabla \phi + \nabla \times \psi$$

$$u_x = \frac{\partial \phi}{\partial x} - \frac{\partial \psi}{\partial z}; u_z = \frac{\partial \phi}{\partial z} - \frac{\partial \psi}{\partial x} \quad (14)$$

$$u_x \sim A \sin(kx - \omega t) \dots$$

$$u_z \sim A \cos(kx - \omega t) \dots$$

u_x, u_z are $\frac{\pi}{2}$ out of phase

Updated by: Sami Alsaadan

Sources: April 13, 2008 by Patricia Gregg. April 8, 2008 lecture.

“An Introduction to Seismology, Earthquakes, And Earth Structure” by Stein & Wyssession (2007).